

# Constructing acoustic timefronts using random matrix theory

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*Europhys. Lett.* **97**, 34002 (2012).

## Collaboration

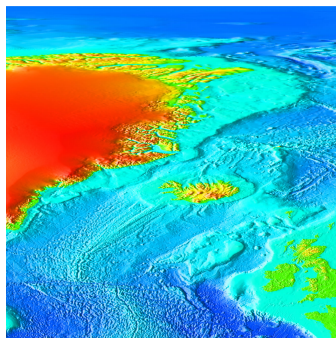
Katherine C. Hegewisch, Ph. D. thesis

# Today's Thread of Logic

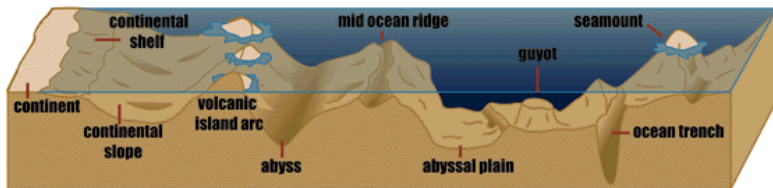
- 1) General considerations and experiments
  - The ocean
    - wave guide with disorder
  - Long range acoustic experiments
    - acoustic timefronts
- 2) Long range propagation models<sup>1</sup>
  - Wave equation
  - One way approximations
  - Paraxial optical approximations
  - Confinement and internal waves
- 3) Introducing Random Matrix Theory
  - Modes and mixing
  - Unitary propagation
  - Constructing acoustic time fronts
- 4) Concluding remarks

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<sup>1</sup>Reviews: M. G. Brown et al., *J. Acoust. Soc. Amer.* **113** (5), 2533 (2003); F. J. Beron-Vera et al., *J. Acoust. Soc. Amer.* **114** (3), 1226 (2003); M. G. Brown and S. Tomsovic, in M. Wright and R. Weaver, editors, *New directions in linear acoustics and vibration: quantum chaos, random matrix theory, and complexity*, CUP, 2010



### Features of the Ocean Floor



Ocean depth: 0-10 km (maximum)

# Considerations

## 1) Short range ocean acoustics

- On continental shelves or inland seas
- Typical ranges of tens of km at most
- Frequencies up to a few kHz ( $c_0 = 1.5 \text{ km/s}$ )
- Surface reflections
- Dissipation quite important, especially bottom interactions

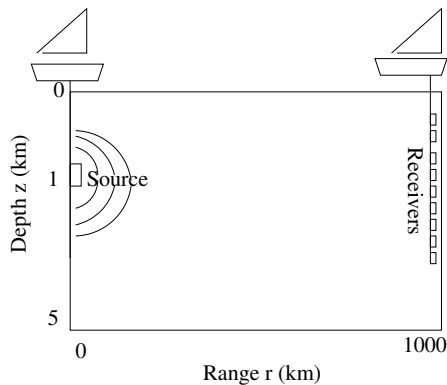
## 2) Long range ocean acoustics

- Align with abyssal plain
- Up to thousands of kilometers
- Lower frequencies - 25 Hz to 250 Hz

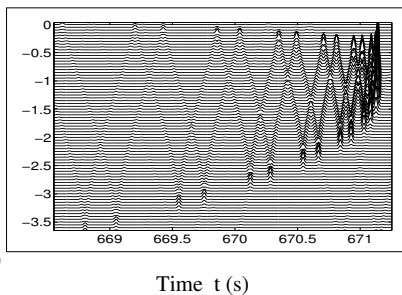
### & Wave guide

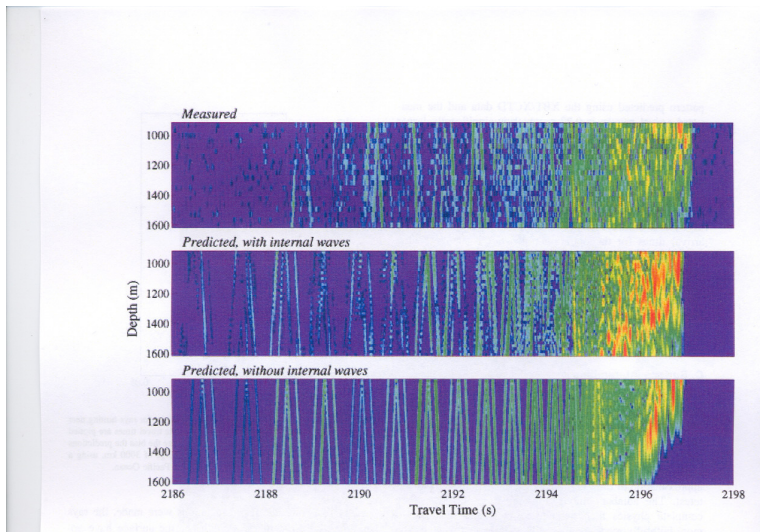
- Warm surface waters
- Constant cold, high pressure waters at depths
- Surface reflections, dissipation and bottom interactions largely avoided

# Measurements



Traces in time at receivers





P. F. Worcester et al., *J. Acoust. Soc. Amer.* **105**, 3185 (1999)

J. A. Colosi et al., *J. Acoust. Soc. Amer.* **105**, 3202 (1999)

# Some quantities of interest

- Mean ocean temperature
- Time front bias, wander, and spread - frequency dependence, range-dependence
- Intensity statistics
- Power distribution and infill
- Questions that have been asked:
  - are ray methods applicable? if so where?
  - how do they relate to mode methods of analysis?

# Wave equation

- 1) A standard approach consists of beginning with the Helmholtz equation:

$$0 = \nabla^2 u(\mathbf{r}; \omega) + \omega^2 c^{-2}(\mathbf{r}) u(\mathbf{r}; \omega)$$

$$\omega = \text{angular frequency}$$

$$c(\mathbf{r}) = \text{position-dependent sound speed}$$

- 2) The time-dependent wave equation solutions

$$\frac{1}{c^2(\mathbf{r})} \frac{\partial^2 \phi(\mathbf{r}; t)}{\partial t^2} = \nabla^2 \phi(\mathbf{r}; t)$$

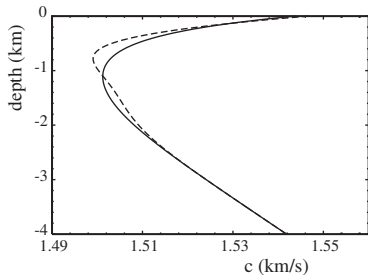
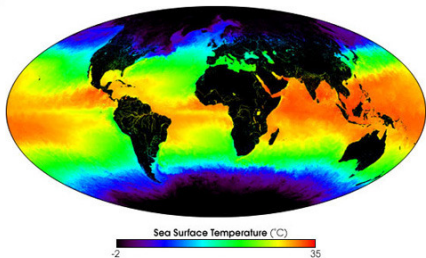
are built as weighted superpositions of eigenstates

$$\phi(\mathbf{r}; t) = \int_{-\infty}^{\infty} d\omega \rho(\omega) e^{-i\omega t} u(\mathbf{r}; \omega)$$

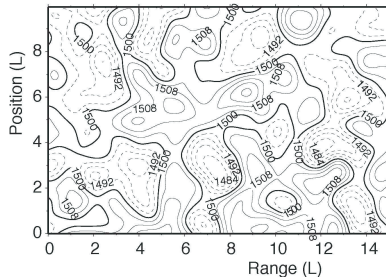
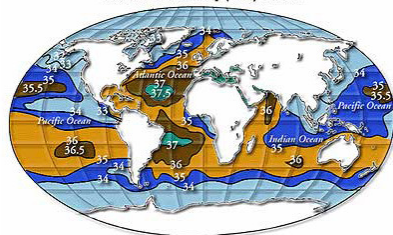
- 3) Boundary conditions:  
determined by ocean surface, bottom, and acoustic source



# Wave guide

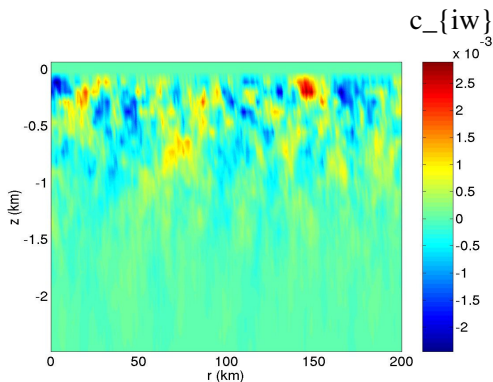


Sea Surface Salinity (SSS) values



# Buoyancy and temperature: internal waves

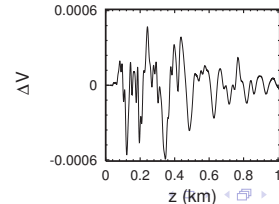
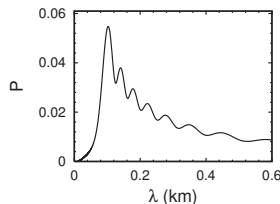
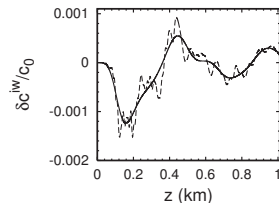
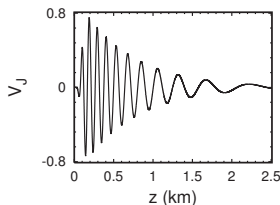
- Vertically displaced water undergoes restoring force
- Strongest force where temperature gradient is strongest
- In mid-latitudes, effect concentrated near surface
- Fluctuation scales range from meters to 100 km
- Vary on minutes to hours time scale
- Responsible for multiple scattering or wave chaos



# Internal waves

Buoyancy modes:

$$\frac{\delta c_{iw}}{c_0} = \sum_{j=1}^{J_{max}} \sum_{k_r} e_{j,k_r} \exp\left(-\frac{3z}{2B}\right) \sin(j\pi\xi(z))$$



# One way approximations

- Range can be used as the time-like variable if there is no backscattering
- In a semiclassical analysis, this leads to a Hamiltonian with a square root
- The quantized version is analogous to a Klein-Gordon equation
- Not a sufficient simplification, consider that the internal waves also can only scatter with small angle changes

# Paraxial approximations - Tappert 1974

A good ansatz for forward-motion and small angle scattering is

$$u(\mathbf{r}; \omega) = \Psi(z, \rho; \omega) \frac{e^{ik_0(\omega)\rho}}{\sqrt{\rho}}$$

Using Helmholtz and dropping small terms gives the parabolic equation

$$\frac{i}{k_0} \frac{\partial}{\partial \rho} \Psi(z, \rho; \omega) = -\frac{1}{2k_0^2} \frac{\partial^2}{\partial z^2} \Psi(z, \rho; \omega) + V(z, \rho) \Psi(z, \rho; \omega)$$

with  $c(z, \rho) = c_0 + \delta c(z, \rho)$  and  $\delta c(z, \rho) \ll c_0$ , the potential is

$$V(z, \rho) = \frac{1}{2} \left( 1 - \left( \frac{c_0}{c(z, \rho)} \right)^2 \right) \approx \frac{\delta c(z, \rho)}{c_0}$$

Notes:

- $\rho \rightarrow t$  and  $k_0 \rightarrow \hbar^{-1}$  gives the Schrödinger equation
- refraction naturally both range and depth dependent

# Introducing random matrix theory for propagation

& So how does one go about constructing a random matrix theory for the propagation of ocean acoustic waves?

- Let's only consider the simplest problem, i.e. that of long range propagation
  - low, fixed frequency (Helmholtz to begin) - will use 75 Hz ahead
  - no losses or dissipation
  - no surface or bottom interactions or absorption
  - no horizontal, out-of-vertical plane scattering
- There is a great deal of deterministic propagation that must be taken into account
- The internal waves create multiple scattering, but have non-zero correlation lengths and are a weak perturbation
- The time-dependent Schrödinger equation leads to unitary propagation

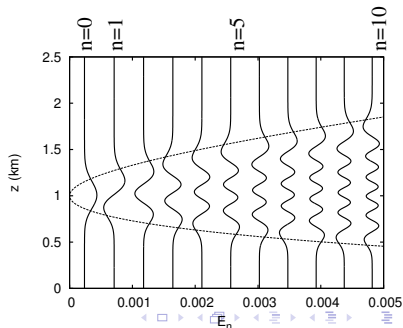
# Modes

Mode picture of propagation (Dozier, Tappert, 1978)

- Modes  $\psi_m$ , energies  $E_m$  of unperturbed waveguide  $V_0$

- Can also be defined adiabatically to account for mesoscale structure
- Somewhere around the 60<sup>th</sup> mode, they begin to hit the surface - will ignore that
- They give a complete representation for the full propagation of the waves

$$-\frac{1}{2k_0^2} \frac{d^2 \psi_m}{dz^2} + V_0(z) \psi_m = E_m \psi_m$$



# Mixing

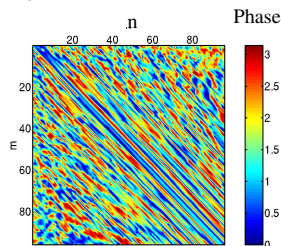
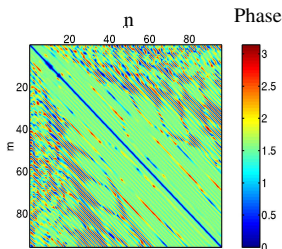
- Unitary propagator coupling coefficients

$$U_{m,n}(\rho; 0) = \int dz \psi_m^*(z) U(\rho; 0) \psi_n(z)$$

gives probability amplitude of mode transition  $n \rightarrow m$

- Unperturbed propagation:  $U_{mn}(\rho; 0) = \Lambda_{mn} = e^{-ik_0 E_n \rho} \delta_{nm}$
- Perturbed propagation: amplitude/phase deviations

$U' = \Lambda^{-1/2} U \Lambda^{-1/2}$  plotted





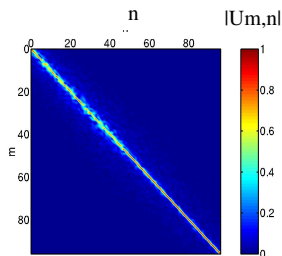
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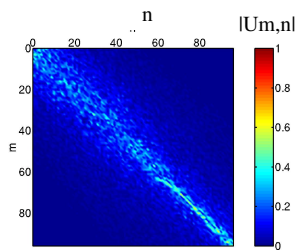
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Propagation to 50 km



Propagation to 1000 km

# Unitary propagation with random matrices

Using building blocks for  $\rho = 50$  km (similar to Perez et al, 2007, for quasi-1D electronic conductors) and a Cayley transformation (the matrix  $A$  Hermitian) for unitarity

$$U = \Lambda^{1/2}(I + i\epsilon A)^{-1}(I - i\epsilon A)\Lambda^{1/2}$$

- Unperturbed result:  $\Lambda_{mn} = e^{-ik_0 E_n \rho} \delta_{mn}$
- Internal wave effects:

$$A_{mn}(k) = \frac{\sigma_{A_{mn}}(k)}{2} z_{mn} \quad \text{i.e. } z_{mn}(k) \text{ perfectly correlated}$$

$$z_{m,n} = \begin{cases} \frac{G(0,1) + iG(0,1)}{\sqrt{2}} & \text{for } n \neq m \\ G(0,1) & \text{for } n = m \end{cases}$$

Long range propagation for  $\rho = 50 \times N$  km

$$U(\rho; 0) = \prod_{i=1}^N U_i(\rho = 50)$$

# Perturbation theory

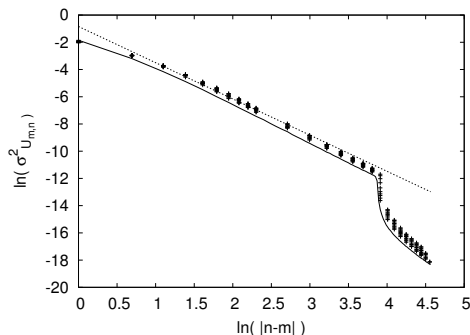
- Using range-dependent (time-dependent) perturbation theory

$$A = \frac{k_0}{2} \int_0^{\rho=50 \text{ km}} d\rho \hat{V}_I$$

where  $\hat{V}_I$  is the operator corresponding to  $\delta c(\mathbf{r})/c_0$  in the interaction picture

- The expectation value of squares of  $A$  matrix elements (variance) can be derived with internal wave formulation of Brown and Colosi, 1998
- They depend almost exclusively on the index difference  $|n - m|$ , i.e.  $A$  is banded with a width depending on  $|n - m|$

- Variance at 75 Hz for 50 km



Approximate fit

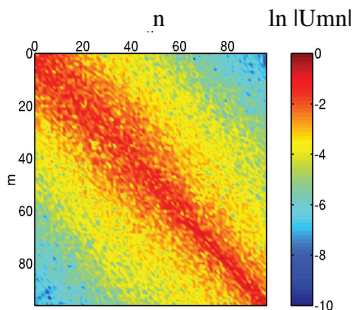
$$\sigma_{U_{m,n}}^2 \approx |n - m|^{-2.6}$$

– Non-unitary, but similar ensemble exhibits localization and superdiffusion (Mirlin et al., 1996)

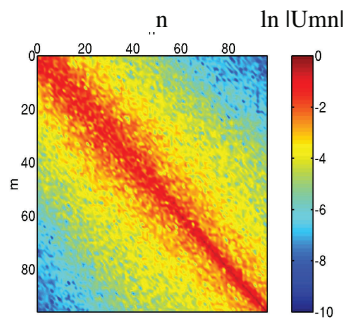
(line=pert. theory, plusses=simulations, dotted line=approx fit)

# Paraxial propagation vs random matrix propagation

- Samples at 75 Hz



(a) wave eqn to 1000 km

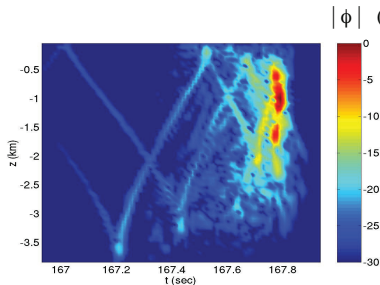


(b) RMT model to 1000 km

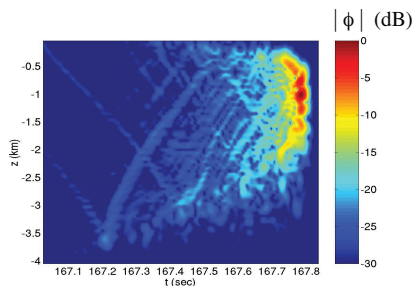
# Acoustic time fronts

- A sample timefront ( $k = \omega/c_0$ )

$$\phi(z, \rho, t) = \frac{1}{\sqrt{2\pi\sigma_k^2}} \int e^{ikc_0(t-\rho/c_0)} u_k(z, \rho) \exp\left[-\frac{(k-k_0)^2}{2\sigma_k^2}\right] dk$$



(a) wave eqn to 250 km

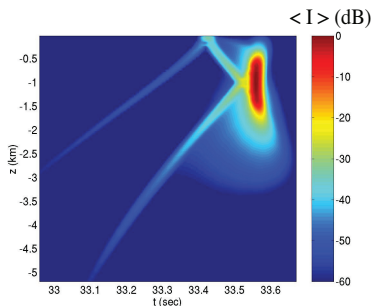


(b) RMT model to 250 km

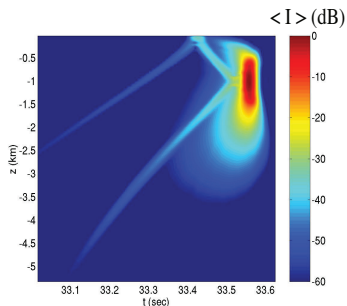
# Averaged timefronts

- Average timefront intensity

$$\langle I \rangle = \frac{1}{N} \sum_{i=1}^N |\phi(z, \rho, t)|^2$$



(a) wave eqn to 50 km

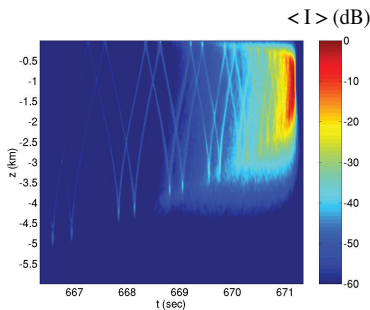


(b) RMT model for 50 km

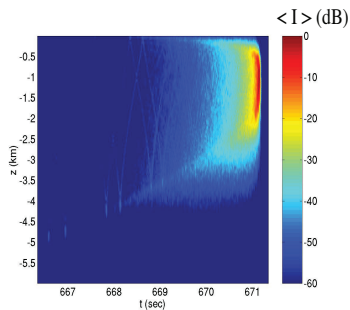
# Averaged timefronts

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(a) wave eqn to 1000 km

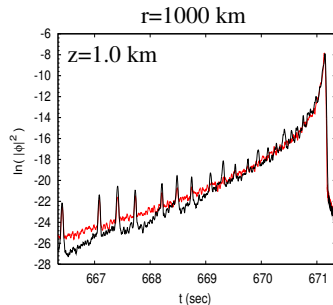
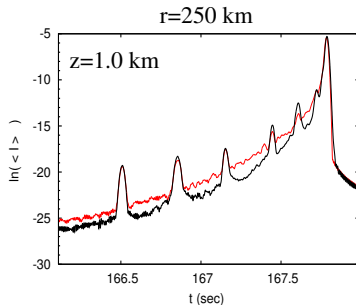


(b) RMT model for 1000 km



# The branches are there

Decay in time of  $\langle I \rangle$  along sound axis



The branches are just enough slightly weaker than in our RMT they are not seen as clearly on the previous slide

# Summary

- The minimum information which must be captured by random matrix ensembles is: i) unitarity; ii) a mean traveling phase for each mode; and iii) a variance decay rate with  $|n - m|$  and  $k$
- Items left out: i) neighboring matrix element correlations; ii) building block correlations; iii)  $k$ -dependence of  $A$  matrix elements; and iv) other ???
- Nevertheless, the ensemble goes a long way to capturing the statistical properties of experimental data
- It would be interesting to: i) investigate the general properties of power-law banded random unitary matrices; ii) understand what other information might yield more faithful RMT propagation (especially  $k$ -dependence of  $z$  matrix element correlations); and iii) learn how to apply RMT to a much broader class of ocean acoustic problems